

Passive Imaging & Monitoring in Wave Physics

New techniques in seismic tomography and joint inversion

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Outline

- Joint body & surface wave traveltime tomography for 3-D Vp, Vs, and Vp/Vs models: methodology and application to NE Tibet
- Direct inversion of 3-D azimuthal and radial anisotropy from surface wave traveltime data: methodology and application to SE Tibet

classic datasets + improved techniques → reliable and useful models

Differences in tomographic models from different datasetsbut we only have one true model!

SE Tibet



Huang et al. (2015)

Different data: different constraints on the model Joint seismic inversion: quest for the true model



Joint inversion: avoid non-uniqueness of inversion using single dataset and get more reliable models

Why joint inversion using body wave + surface wave traveltimes?

Complementary strengths

• 1. Different depth and horizontal sensitivities



West et al. (2004, GRL)

Why joint inversion using body wave + surface wave traveltimes?

• 2. Different model parameter sensitivity



Background: surface wave tomography



Fang, Yao, Zhang et al. (2015, GJI)

Direct inversion of 3-D Vs model from dispersion data with period dependent ray tracing



code available at https://github.com/HongjianFang/DSurfTomo

Applications of DSurfTomo

- Regional scale (a few hundred to thousand km→ crustal structure): Tibetan plateau, SW Tibet, SE China and Taiwan Straight, eastern China ...
- Local scale (~ ten to hundred km→shallow crust): Tanlu fault zone, Hefei City, Jinan City, Taipei Basin, Binchuan Basin in SW China,
- Exploration scale (several km→near surface): shale gas production field, gas storage place,

Joint body & surface wave traveltime tomography for 3-D Vp and Vs models

Body wave Traveltime tomo

Direct Surface wave Traveltime tomo

$$\begin{bmatrix} \mathbf{G}_{V_p}^{SW} & \mathbf{G}_{V_s}^{SW} \end{bmatrix} \begin{bmatrix} \Delta \mathbf{m}_p \\ \Delta \mathbf{m}_s \end{bmatrix} = \mathbf{d}^{SW}$$

 $\begin{bmatrix} \mathbf{G}_{H}^{T_{p}} & \mathbf{G}_{Vp}^{T_{p}} & \mathbf{0} \\ \mathbf{G}_{H}^{T_{s}} & \mathbf{0} & \mathbf{G}_{Vs}^{T_{s}} \end{bmatrix} \begin{bmatrix} \Delta \mathbf{H} \\ \Delta \mathbf{m}_{p} \\ \Delta \mathbf{m}_{p} \end{bmatrix} = \begin{bmatrix} \mathbf{d}^{T_{p}} \\ \mathbf{d}^{Ts} \end{bmatrix}$

period-dependent SW ray tracing (Fang, Yao, Zhang et al. 2015, GJI.)

Joint body & surface waves tomo

$$\begin{bmatrix} \mathbf{G}_{H}^{T_{p}} & \mathbf{G}_{Vp}^{T_{p}} & \mathbf{0} \\ \mathbf{G}_{H}^{T_{s}} & \mathbf{0} & \mathbf{G}_{Vs}^{T_{s}} \\ \mathbf{0} & \alpha \mathbf{G}_{Vp}^{SW} & \alpha \mathbf{G}_{Vs}^{SW} \end{bmatrix} \begin{bmatrix} \Delta \mathbf{H} \\ \Delta \mathbf{m}_{p} \\ \Delta \mathbf{m}_{s} \end{bmatrix} = \begin{bmatrix} \mathbf{d}^{T_{p}} \\ \mathbf{d}^{Ts} \\ \alpha \mathbf{d}^{SW} \end{bmatrix}$$

(Fang, Zhang, Yao et al., JGR 2016)

Joint body & surface wave traveltime tomography for 3-D Vp and Vs models



 λ_1, λ_2 : 3-D spatial smoothing λ_3, κ : Vp/Vs ratio prior constraints $(\mathbf{m}_p + \Delta \mathbf{m}_p) = \kappa(\mathbf{m}_s + \Delta \mathbf{m}_s)$

(Fang, Zhang, Yao et al., JGR 2016)

Joint Inversion for 3D Vp/Vs ratio

Fang, Yao, Zhang et al. (GJI, 2019)



Add smoothing and damping directly to Vp/Vs to stabilize the inversion of 3-D Vp/Vs → lithology, partial melting?

Application to NE Tibetan Plateau



Body wave traveltime data: P picks: ~ 300,000 S picks: ~ 290,000 (from Shunping Pei) Surface wave traveltime data: Rayleigh wave phase velocity dispersion: ~ 51,000 (10 to 41 s) (from Hongyi Li)

Previous noise tomography results

prominent low-velocity zone in the middle crust of Qiangtang and Songpan-Ganze Terranes; no clear evidence of northeastward crustal flow to the Qilian Orogen



Joint inversion: synthetic tests



Vp

Vs





Output model for Vp/Vs



Horizontal slices at different depths



Horizontal slices at different depths



Vp/Vs model at different depths







1.7 1.8 Vp/Vs

Fang et al. (in prep)

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Azimuthal Anisotropy

HTI medium: horizontal sym. axis



Radial Anisotropy

VTI medium: vertical sym. axis



Causes for seismic anisotropy: cracks, layering, shape or lattice preferred orientation

Methods for seismic anisotropy: local S-wave or XKS splitting, receiver functions (Pms), body wave traveltimes, surface wave dispersion, etc

Classical representation of Rayleigh-wave and shear-wave velocity azimuthal anisotropy



Shear-wave velocity:

$$\hat{\beta}_{SV} \approx V_{SV} \left(1 + \frac{G_c}{2L} \cos 2\psi + \frac{G_s}{2L} \sin 2\psi \right)$$
$$= V_{SV} \left[1 + \Lambda_{SV} \cos 2(\psi - \phi_F) \right]$$

Smith & Dahlen (1973) Montagner & Nataf (1986)

1-D depth sensitivity kernels

Normal mode / surface wave mode theory



Weak azimuthal anisotropy : $V_{SV} \approx V_{SV} \left(1 + \frac{G_c}{2L} \cos 2\psi + \frac{G_s}{2L} \sin 2\psi\right)$

Traditional two-step inversion for Vs azimuthal anisotropy from Rayleigh waves



Traditional two-step inversion for Vs azimuthal anisotropy from Rayleigh waves



Direct inversion for 3-D Vs azimuthal anisotropy based on raytracing from dispersion data



Step 1. All dispersion data \rightarrow 3-D isotropic Vs



Fang, Yao, et al. (2015, GJI)

Step 2. All traveltime residual data \rightarrow 3-D isotropic and azimuthally anisotropic Vs

Based on the 3-D ref. model from Step 1

$$\delta t_{i}(\omega) = t_{i}^{obs}(\omega) - t_{i}^{ref}(\omega) \approx \sum_{k=1}^{K} \frac{-\mu_{ik}}{(c_{0}^{k}(\omega))^{2}} \left(\delta c_{k}(\omega) + a_{1}^{k}(\omega) \cos 2\psi + a_{2}^{k}(\omega) \sin 2\psi \right)$$

$$= \sum_{k=1}^{K} \frac{-\mu_{ik}}{(c_{0}^{k}(\omega))^{2}} \left(\delta c_{k}^{ETI}(\omega) + \delta c_{k}^{AA}(\omega, \psi) \right).$$

$$\delta c_k^{ETI}(\omega) = \int_0^\infty \left(\frac{\partial c_k(\omega)}{\partial \alpha_k(z)} \delta \alpha_k(z) + \frac{\partial c_k(\omega)}{\partial \beta_k(z)} \delta \beta_k(z) + \frac{\partial c_k(\omega)}{\partial \rho_k(z)} \delta \rho_k(z) \right) dz,$$

$$\begin{split} \delta c_k^{AA}(\omega) &= \int_0^\infty \left\{ \left(B_c^k \frac{\partial c_k(\omega)}{\partial A_k} + G_c^k \frac{\partial c_k(\omega)}{\partial L_k} \right) cos 2\psi + \left(B_s^k \frac{\partial c_k(\omega)}{\partial A_k} + G_s^k \frac{\partial c_k(\omega)}{\partial L_k} \right) sin 2\psi + (B_c^k \frac{\partial c_k(\omega)}{\partial A_k} + G_s^k \frac{\partial c_k(\omega)}{\partial L_k} \right) sin 2\psi + (B_c^k \frac{\partial c_k(\omega)}{\partial A_k} + G_s^k \frac{\partial c_k(\omega)}{\partial L_k} + G_s^k \frac{\partial c_k(\omega)}{\partial L_k} \right) sin 2\psi + (B_c^k \frac{\partial c_k(\omega)}{\partial A_k} + G_s^k \frac{\partial c_k(\omega)}{\partial L_k} + G_s^k \frac{\partial c_k(\omega)}{\partial L_k} + G_s^k \frac{\partial c_k(\omega)}{\partial L_k} \right) sin 2\psi + (B_c^k \frac{\partial c_k(\omega)}{\partial A_k} + G_s^k \frac{\partial c_k(\omega)}{\partial L_k} \right) sin 2\psi + (B_c^k \frac{\partial c_k(\omega)}{\partial A_k} + G_s^k \frac{\partial c_k(\omega)}{\partial L_k} + G$$

Step 2. All traveltime residual data \rightarrow **3-D isotropic and azimuthally anisotropic Vs** some simplifications and finally ... $\delta t_i(\omega) = \sum_{k=1}^K \frac{-\mu_{ik}}{(c_k^k(\omega))^2} \sum_{j=1}^J \left\{ \left[\left(\int_{z_j}^{z_{j+1}} \frac{\partial c_k(\omega)}{\partial \alpha_k(z)} \, dz \right) R_\alpha(z_j) + \left(\int_{z_j}^{z_{j+1}} \frac{\partial c_k(\omega)}{\partial \beta_k(z)} \, dz \right) + \right] \right\}$ $\left(\int_{z_{j}}^{z_{j+1}} \frac{\partial c_{k}(\omega)}{\partial \rho_{k}(z)} dz\right) R_{\rho}(z_{j}) \left[\delta\beta_{k}(z_{j}) + \eta_{ik} \left[\left(\int_{z_{j}}^{z_{j+1}} \frac{\partial c_{k}(\omega)}{\partial A_{k}(z)} dz\right) A_{k}(z_{j}) + \left(\int_{z_{j}}^{z_{j+1}} \frac{\partial c_{k}(\omega)}{\partial L_{k}(z)} dz\right) L_{k}(z_{j})\right] \frac{G_{c}^{\kappa}(z_{j})}{L_{k}(z_{j})} + \frac{G_{c}^{\kappa}(z_{j})}{G_{c}^{\kappa}(z_{j})} dz$ $\xi_{ik} \left[\left(\int_{z_j}^{z_{j+1}} \frac{\partial c_k(\omega)}{\partial A_k(z)} \, dz \right) A_k(z_j) + \left(\int_{z_j}^{z_{j+1}} \frac{\partial c_k(\omega)}{\partial L_k(z)} \, dz \right) L_k(z_j) \right] \frac{G_s^k(z_j)}{L_k(z_j)} \right\},$ → d = G m $\mathbf{m} = \left[\Delta \beta_1(z_j) \dots \Delta \beta_1(z_J) \dots \Delta \beta_K(z_J) \frac{G_c^1(z_1)}{L_1(z_1)} \dots \frac{G_c^1(z_J)}{L_1(z_I)} \dots \frac{G_c^K(z_J)}{L_K(z_J)} \frac{G_s^1(z_1)}{L_1(z_1)} \dots \frac{G_s^1(z_J)}{L_1(z_I)} \dots \frac{G_s^K(z_J)}{L_K(z_J)} \right]^T$ Inversion matrix: $\begin{array}{ccc} \mathbf{U}_{iso} & \mathbf{U}_{AA} \\ \lambda_1 \mathbf{L}_{iso} & \mathbf{0} \\ \mathbf{0} & \lambda_2 \mathbf{L}_{idd} \end{array} \begin{bmatrix} m_{iso} \\ m_{AA} \end{bmatrix} = \begin{bmatrix} a \\ \mathbf{0} \\ \mathbf{0} \end{bmatrix}$

Synthetic examples



DAzimSurfTomo: Application to SE Tibet





Vs model in the crust and uppermost mantle



Liu, Yao, Yang et al. (JGR, 2019)

Comparison of 2-D phase v maps



1° x 1° checkerboard model







(Chang et al, 2016)

Apparent differences in crust and uppermost mantle azim. aniso.



Liu, Yao, Yang et al. (JGR, 2019)

Direct inversion for 3-D Vsh and Vsv

Rayleigh
$$\left(\begin{array}{c} \Delta \mathbf{T}_{R} \\ \Delta \mathbf{T}_{L} \end{array} \right) = \left(\begin{array}{c} \mathbf{G}_{sv} & \mathbf{0} \\ \mathbf{0} & \mathbf{G}_{sh} \end{array} \right) \left(\begin{array}{c} \Delta \mathbf{V}_{sv} \\ \Delta \mathbf{V}_{sh} \end{array} \right)$$

3-D inversion: spatial smoothing added on ΔV_{sh} & ΔV_{sv}

Radial anisotropy : direct division \rightarrow large uncertainty

Direct inversion for 3-D Vs radial anisotropy

$$\gamma = \frac{V_{sh}}{V_{sv}} \implies \Delta \mathbf{V}_{sh} = \gamma \cdot \Delta \mathbf{V}_{sv} + \mathbf{V}_{sv} \cdot \Delta \gamma$$

$$\begin{pmatrix} \Delta \mathbf{T}_R \\ \Delta \mathbf{T}_L \end{pmatrix} = \begin{pmatrix} \mathbf{G}_{sv} & \mathbf{0} \\ \mathbf{0} & \mathbf{G}_{sh} \end{pmatrix} \begin{pmatrix} \Delta \mathbf{V}_{sv} \\ \Delta \mathbf{V}_{sh} \end{pmatrix}$$

$$\begin{pmatrix} \Delta \mathbf{T}_R \\ \Delta \mathbf{T}_L \end{pmatrix} = \begin{pmatrix} \mathbf{G}_{sv} & \mathbf{0} \\ \gamma \cdot \mathbf{G}_{sh} & \mathbf{V}_{sv} \cdot \mathbf{G}_{sh} \end{pmatrix} \begin{pmatrix} \Delta \mathbf{V}_{sv} \\ \Delta \gamma \end{pmatrix}$$
More stable: spatial smoothing directly added

$$\left(\frac{V_{sh}+V_{sv}}{2}, \frac{2(V_{sh}-V_{sv})}{V_{sh}+V_{sv}}\right) \bigstar \left(V_{sv}\frac{\gamma+1}{2}, \frac{2(\gamma-1)}{\gamma+1}\right)$$

DRadiSurfTomo: application to SE Tibet



Synthetic examples: 3-D radial anisotropy



Direct inversion for radial aniso. ($\Delta V_{sv}, \Delta \gamma$)



Hu, Yao, Huang (submitted to JGR)

3-D radial anisotropy in the upper and middle crust

Hu, Yao, Huang (submitted to JGR)



Conclusions

- Developed the joint inversion method of body & surface wave traveltime → more reliable 3-D Vp, Vs, and Vp/Vs models. Future work will include the station-based RFs and ZH data.
- Developed direct inversion methods for 3-D azimuthal and radial anisotropy from surface wave traveltime data. Future work will include body wave traveltimes for 3-D joint anisotropy inversion.

Thank you!

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DSurfTomo package download: https://github.com/HongjianFang/DSurfTomo DAzimSurfTomo package download: https://github.com/Chuanming-Liu/DAzimSurfTomo DBodySurfTomo & DRadiSurfTomo packages in progress